

雲マイクロ物理過程と乱流スカラー輸送の大規模シミュレーション

名古屋大学HPC 計算科学連携研究プロジェクト成果報告シンポジウム

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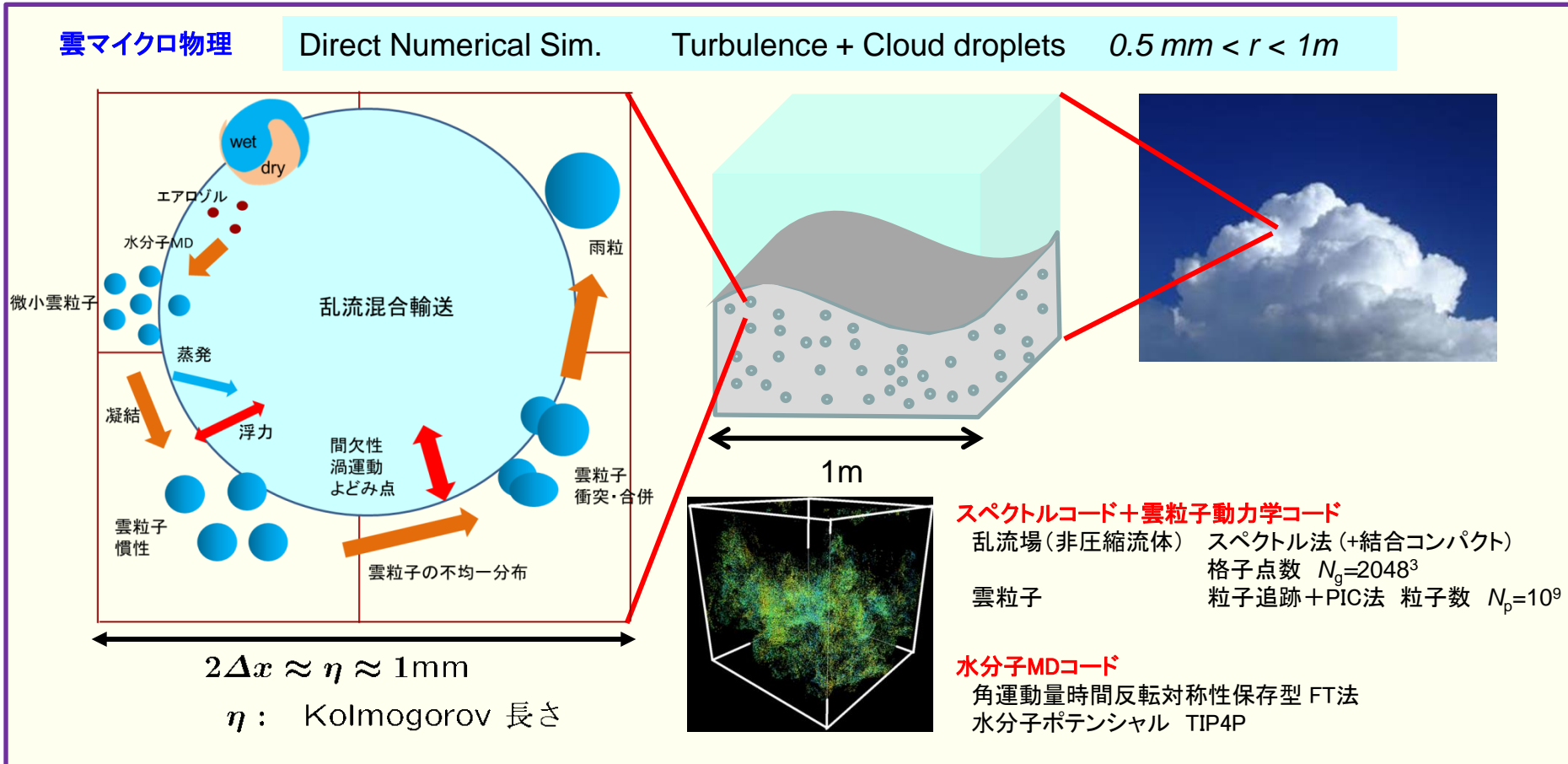
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雲粒子はどのようにして生成され、成長し、輸送・混合され雨粒になるか？

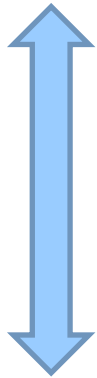
- ・ 雲粒子と乱流による水蒸気や温度の混合・輸送との相互作用におけるマイクロからセミマクロまでのプロセスを物理の基本原則にのっとり丸ごとシミュレーションできるプログラムを開発
- ・ 雲粒子形成のきっかけ、凝結成長、乱流による雲粒子と水蒸気の不均一な空間分布の形成、凝結と衝突による粒径分布の変化、乱流強へのフィードバックなどを調べて、雨粒形成までのプロセスを解明する



Motivation 1

Turbulence Cloud interaction

Turbulence: transport and mixing of momentum, heat and mass
many scales of motion
Eulerian representation



Interaction buoyancy (heat, water vapor), momentum
small scales or all scales, time scale

Cloud: water droplets, radius, mass, heat
particle interaction (collision)
Lagrangian representation

Motivation 2

Development of high performance code of turbulence + particles

Many nodes + Many cores

Eulerian code for continuum

Incompressible fluid

Poisson equation : non-local in space $N^3 \log_2 N$

Scalar equation : local in space

Needs for Accuracy and Less communication

Hybrid scheme

spectral + combined compact difference) for scalar solver

acceleration 30%(Sc=1), 400% (Sc>50)

Lagrangian code for particle

Particle tracking random access to the memory

Relabeling

interpolation

Direct Numerical Simulation of Turbulence and Cloud Droplets

Andfrejczuk et al JAP. 2004

Kumar, Schumacher, Shaw, TCFD 2012

Included

- stratocumulus cloud at about 1500m
- turbulence + buoyancy
- temperature
- water vapor mixing ratio
- water droplets of radius $10\mu\text{m} \sim 20\mu\text{m}$
- condensation, evaporation
- Stokes drag + radius dependent relaxation time + gravity
- Collision

Not included

- Collision of droplets, coagulation
- Nucleation of water droplets
- Rain, ice

Goal

Cloud droplets

- How fast do the droplets grow in mean and distribution ?
- What are key processes ?
- What is the spatial distribution of cloud droplets?

Turbulence

- What is role of turbulence in cloud evolution?
- To what extent is turbulence modified or driven?
- What is intermittency of velocity and scalar in turbulence?

Basic Equations

Turbulence (Eulerian)

Boussinesq approximation

$$\frac{\partial u}{\partial t} + u \cdot \nabla u = -\nabla p + \nu \nabla^2 u + e_z B + f, \quad \nabla \cdot u = 0$$

buoyancy external force

$$\frac{\partial T}{\partial t} + u \cdot \nabla T = \kappa \nabla^2 T + \frac{L}{c_p} C_d$$

$$\frac{\partial q_v}{\partial t} + u \cdot \nabla q_v = \kappa \nabla^2 q_v - C_d$$

condensation, evaporation

$$B = g \left(\frac{T - T_0}{T_0} + \epsilon (q_v - q_{v0}) - q_l \right)$$

High Reynolds number turbulence : Spectral method

Scalar transport : Spectral (or hybrid method)

Andfrejczuk et al JAP. 2004

Kumar, Schumacher, Shaw, TCFD 2012

Cloud droplets (Lagrangian)

$$\frac{d\mathbf{X}_j}{dt} = \mathbf{V}_j(t)$$

$$\frac{d\mathbf{V}_j}{dt} = \frac{1}{\tau_j(t)} (\mathbf{u}(\mathbf{X}_j(t), t) - \mathbf{V}_j(t)) + g\mathbf{e}_3 \quad \text{Stokes approximation}$$

$$R_j(t) \frac{dR_j(t)}{dt} = KS(\mathbf{X}_j(t), t), \quad R_j = \text{droplet radius} \quad \text{Diffusion process}$$

$$C_d(x, t) \equiv \frac{1}{m_{air}} \frac{dm_l(x, t)}{dt} = \frac{4\pi r_l K}{\rho_0(\Delta x)^3} \sum_{k=1}^{N_\Delta} R_k(t) S(\mathbf{X}_k(t), t) \quad \text{Condensation rate}$$

$$S = \frac{q_v}{q_{vs}(T)} - 1, \quad \text{supersaturation rate}$$

$$K^{-1} = \frac{\rho_l R_v T}{D_v e_{sat}(T)} + \frac{\rho_l L}{\kappa_a T} \left(\frac{L}{R_v T} - 1 \right)$$

PIC

Interpolation of velocity and scalar fields at particle position

Redistribution of cloud properties onto grid points

Simulation of cloud evolution

Domain : $(25.6 \text{ cm})^3 \rightarrow (102.4 \text{ cm})^3$

Water droplets : sphere (initially $10 \sim 20 \mu\text{m}$)

Turbulence : homogeneous isotropic steady

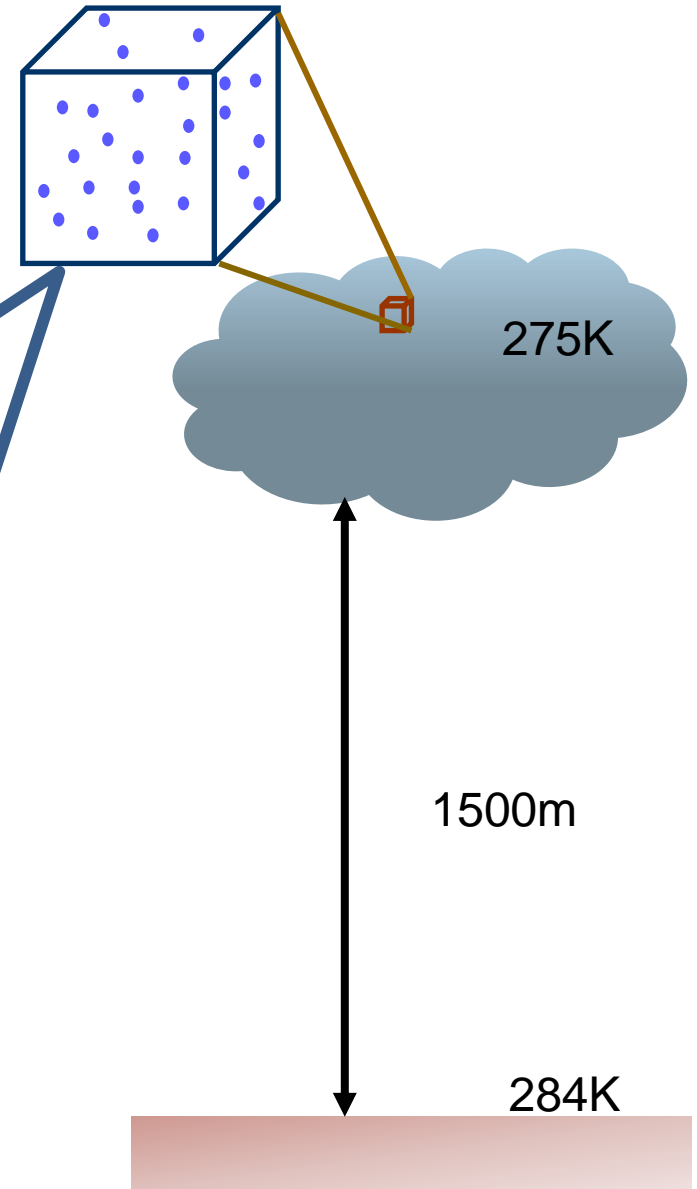
DNS : turbulence + scalar

pseudo spectral method
3DFFT + MPI + Open MP

interpolation for droplets

PIC + TS13 (or **linear**)

4th order Runge-Kutta-Gill

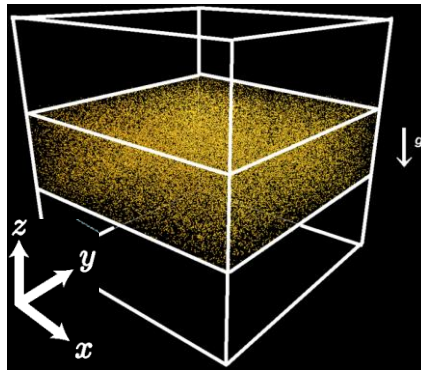


Initial Condition

Turbulence

Temperature fluctuation

Droplets



Water vapor mixing ratio q_v

$$q_v(x, t = 0) = (q_v^{\max} - q_{v0}) \exp(-Az^6) + q_{v0}$$

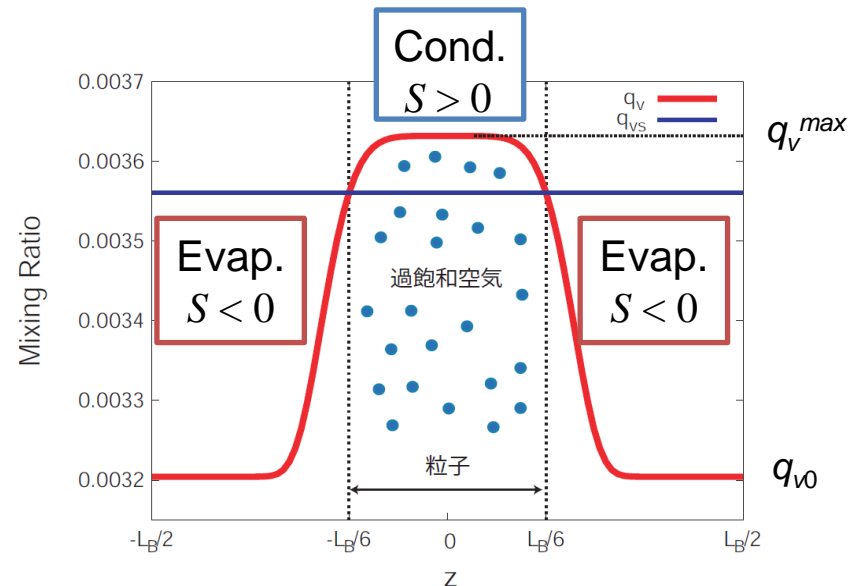
$$q_v^{\max} = 1.02q_{vs} \quad q_{v0} = 0.90q_{vs}$$

$$q_v = q_{vs} \quad \text{at} \quad z = \pm L_B/6$$

Isotropic Steady turbulence

θ : 275 K + zero fluctuation

- Random in space in the range $-L_B/6 \leq z \leq L_B/6$
- No. of droplets : $2^{21} \doteq 2 \times 10^6$
 $2^{27} \doteq 1.3 \times 10^8$
- No. Density of droplets 31 ~ 125/cm³
- Initial radius : 10 ~ 20 μm



Re数、計算領域、格子数、粒子数を増大

Taylor micro scale Re 数: R_λ	92	2.7倍	252
立方体の1辺の長さ: L	25.6cm	4倍	102.4cm
格子点数: N^3	128^3	8倍	1024^3
雲粒子数: N_p	2^{22} 個	2倍	2^{27} 個

Simulations of Decay wet, Decay dry

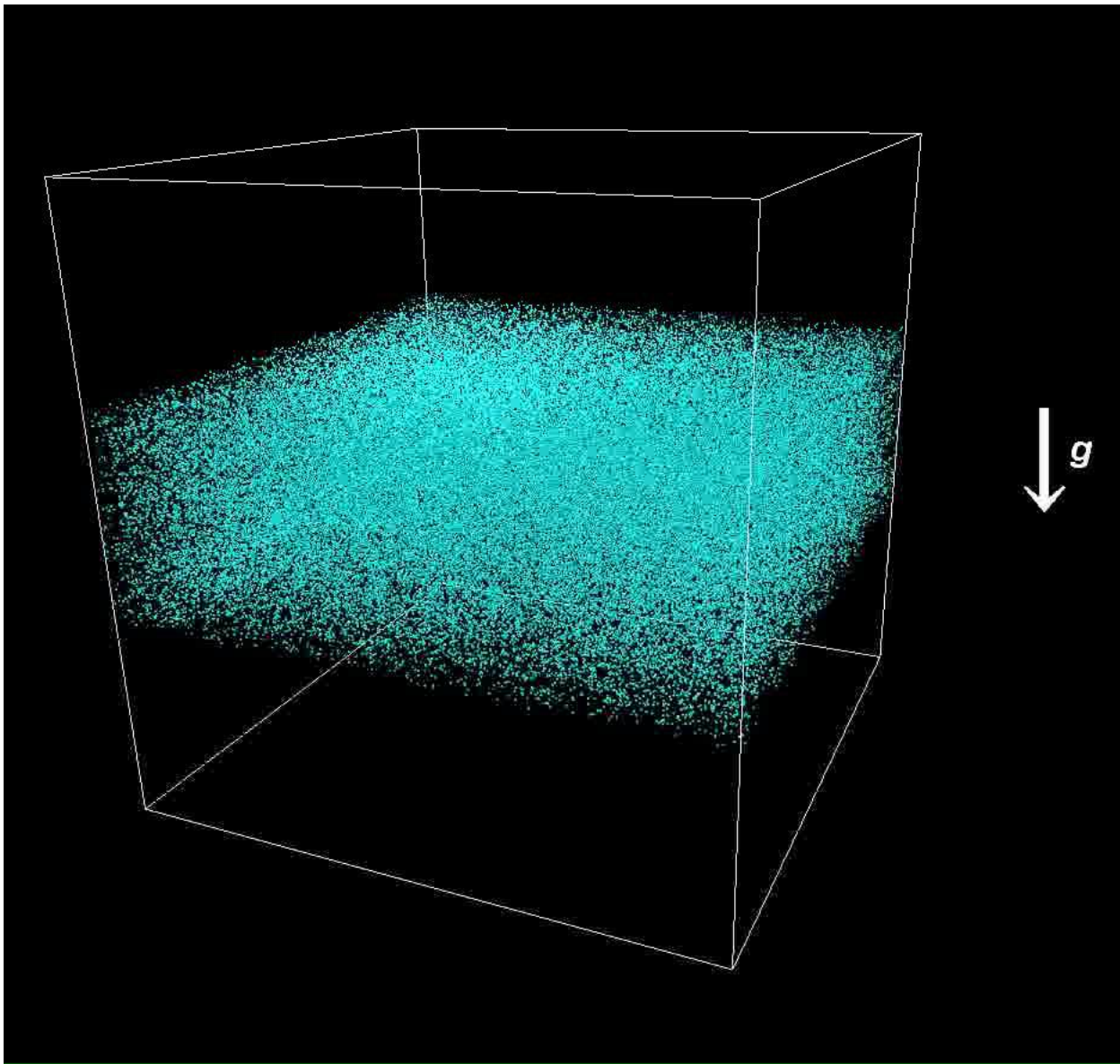
$$\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} = -\frac{1}{\rho_a} \nabla p + \nu_a \nabla^2 \mathbf{u} + B \mathbf{e}_3 + \mathbf{f}$$

	Initially steady	Decay wet	Decay dry
External force F	ON	OFF	OFF
Buoyancy B	OFF	ON	OFF

	A	B	C'	C''	C	D(A)	E(B)
N^3	128^3	128^3	128^3	128^3	1024^3	1024^3	1024^3
N_p	2^{21}	2^{21}	2^{22}	2^{19}	2^{25}	2^{27}	2^{27}
n_p [cm^{-3}]	125	125	250	31	31	125	125
q_v	やせる	ふとる	ふとる	ふとる	ふとる	やせる	ふとる
θ'	0	0	0	0	0.05K	0	0
R_λ	92	92	92	92	252	252	252
$K_{\max}\eta$	1.2	1.2	1.2	1.2	2.1	2.1	2.1
$r(0)$ [μm]	20	20	15	15	15	20	20

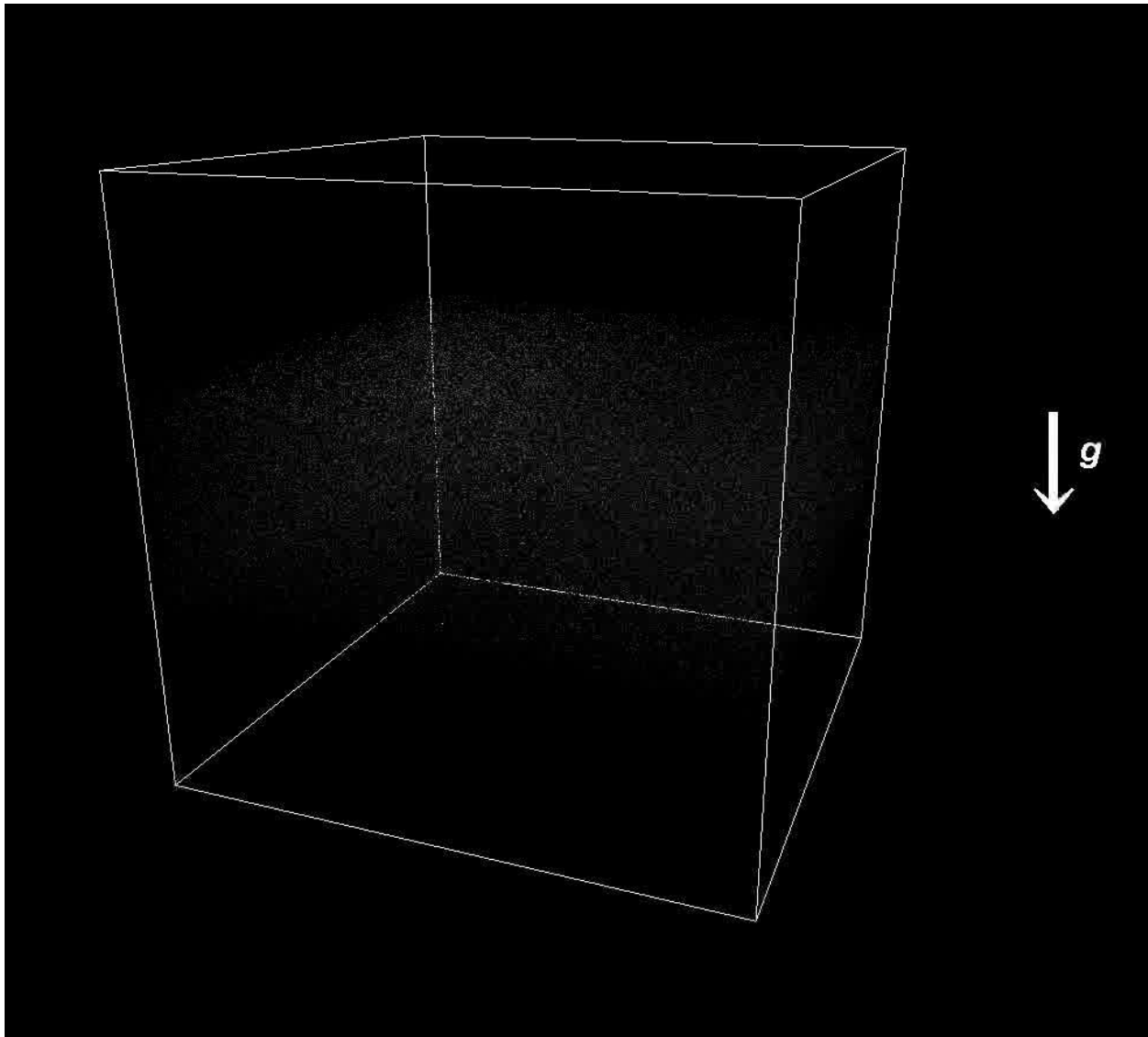
雲粒子 ふとる $M_{vs} \leq M_v + M_l \iff \rho_a \int_V q_{vs} dV \leq \rho_a \int_V q_v(x, t) dV + \frac{4}{3} \pi \rho_l \sum_{j=1}^{N_p} r_j^3$

雲粒子 やせる $M_{vs} \geq M_v + M_l$



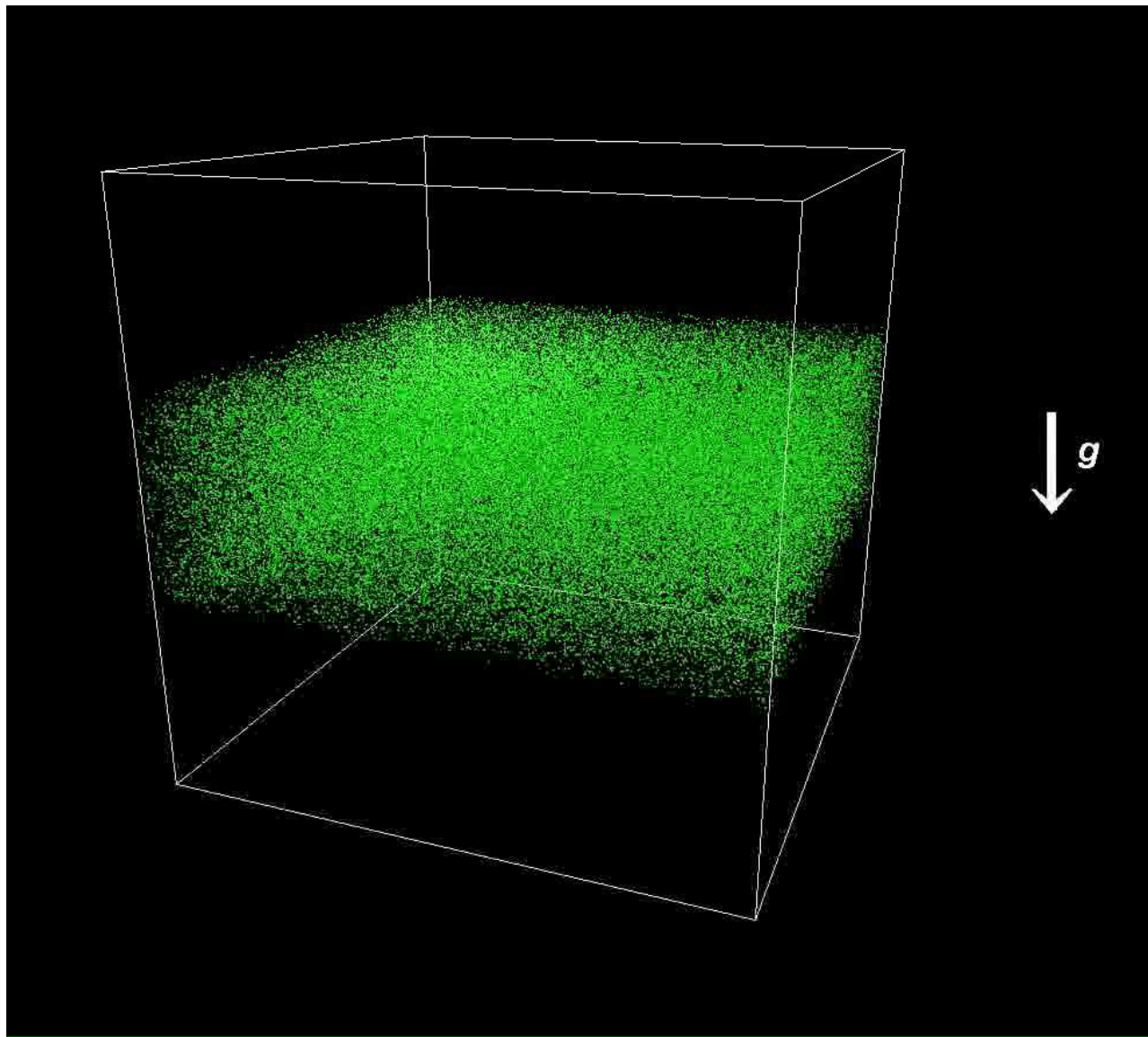
$$R_\lambda = 92$$

- ・境界: 3軸方向周期境界条件
- ・雲粒子の個数: 4×10^6 個
- ・可視化粒子数: 10^5 個
- ・初期雲粒子半径: $15 \mu\text{m}$
- ・粒子の色: 粒子半径
赤 \Rightarrow 半径大、青 \Rightarrow 半径小
- ・0s \sim 4s まで可視化



$$R_\lambda = 252$$

- ・境界: 3軸方向周期境界条件
- ・雲粒子の個数: 1.3×10^8 個
- ・可視化粒子数: 10^5 個
- ・初期雲粒子半径: $20 \mu\text{m}$
- ・ $r > 20 \mu\text{m}$ を黒から白色
- ・0s~6s まで可視化

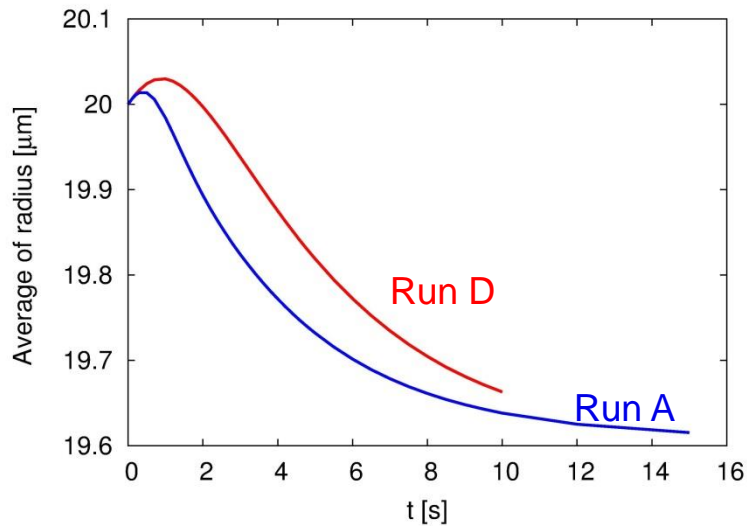


$$R_{\lambda} = 252$$

- ・境界: 3軸方向周期境界条件
- ・雲粒子の個数: 1.3×10^8 個
- ・可視化粒子数: 10^5 個
- ・初期雲粒子半径: $20 \mu\text{m}$
- ・粒子の色: 粒子半径
赤 \Rightarrow 半径大、青 \Rightarrow 半径小
- ・0s \sim 6s まで可視化

Effects of Reynolds number

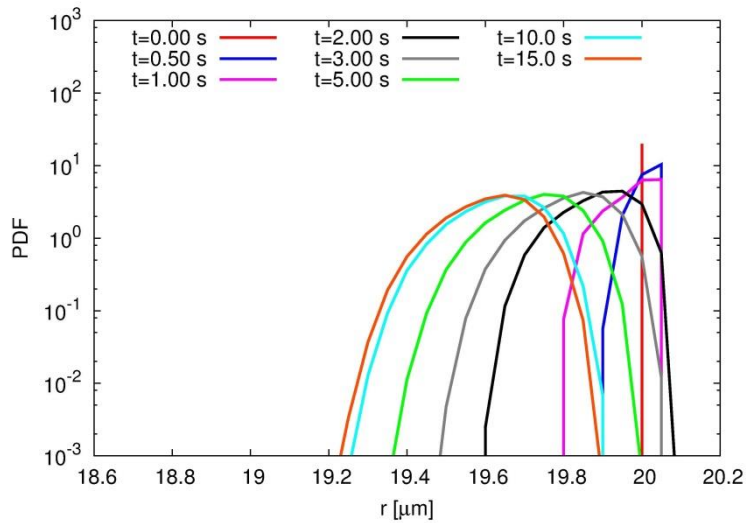
Run A ($R_\lambda = 92$) and D ($R_\lambda = 252$)



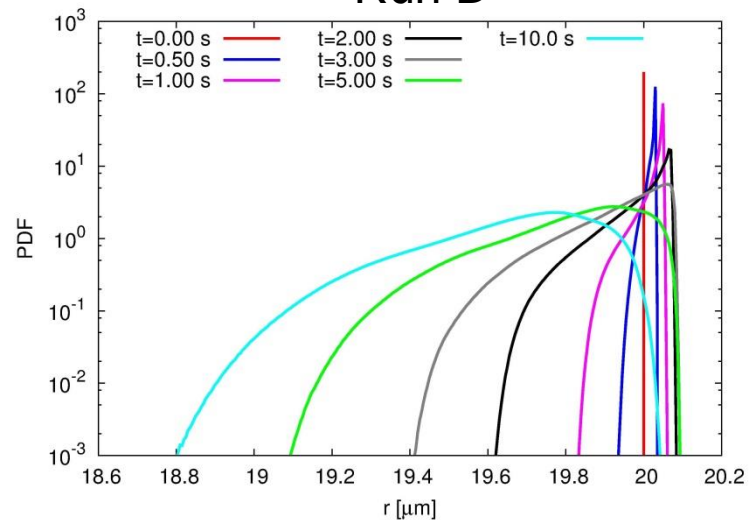
When R_λ increases

- $\langle r \rangle$ → slowly decrease
- PDF → broad

Run A

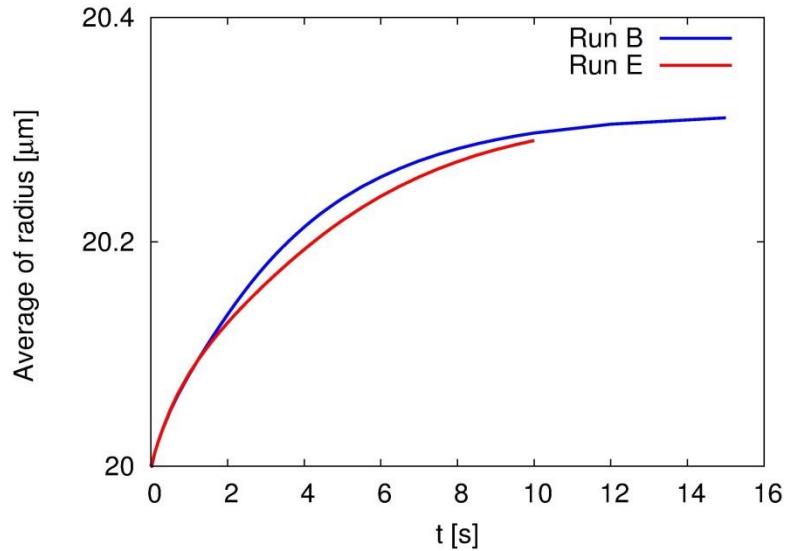


Run D



Effects of Reynolds number

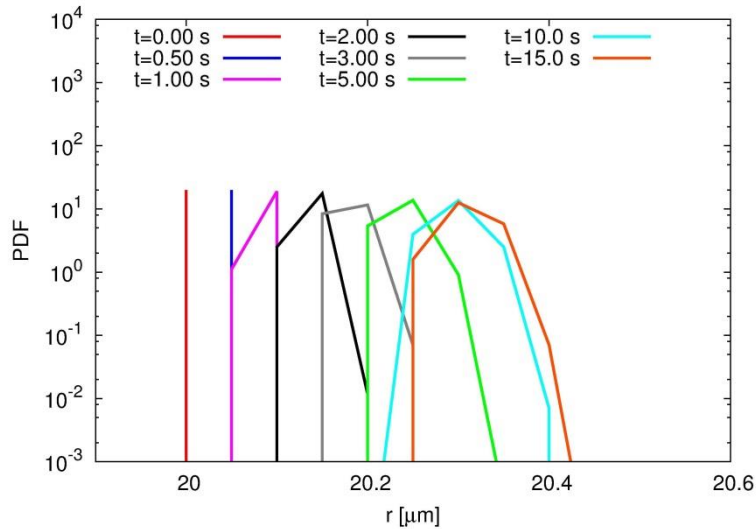
Run B ($R_\lambda=92$) and E ($R_\lambda=252$)



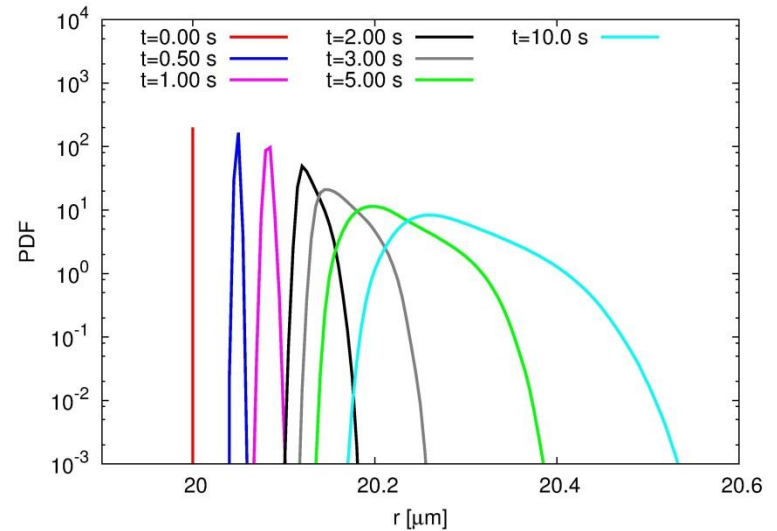
When R_λ increases

$\langle r \rangle$ → comparable
PDF → broad

Run B



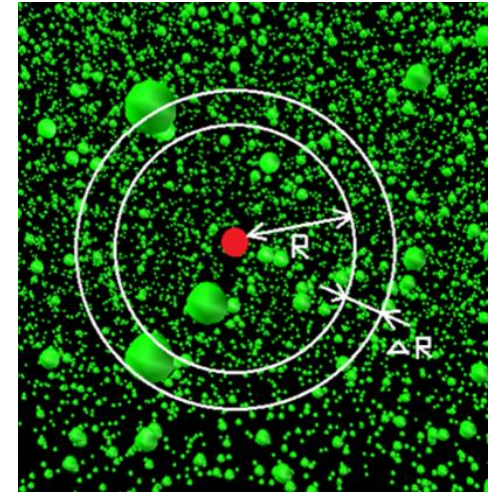
Run E



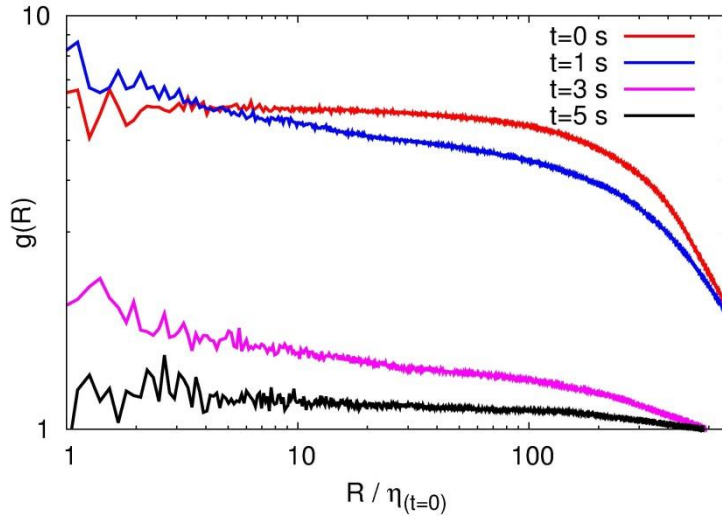
Radial distribution function of cloud droplets

$$g(R) = \frac{G(R)}{4\pi R^2 n_d \Delta R}$$

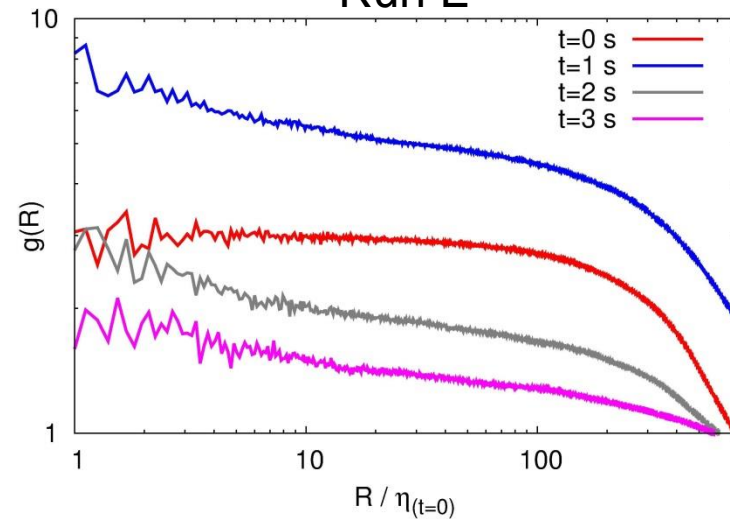
$G(R)\Delta R$: No. of particles within spherical shell between R and ΔR
 n_d : Average number density



Run D

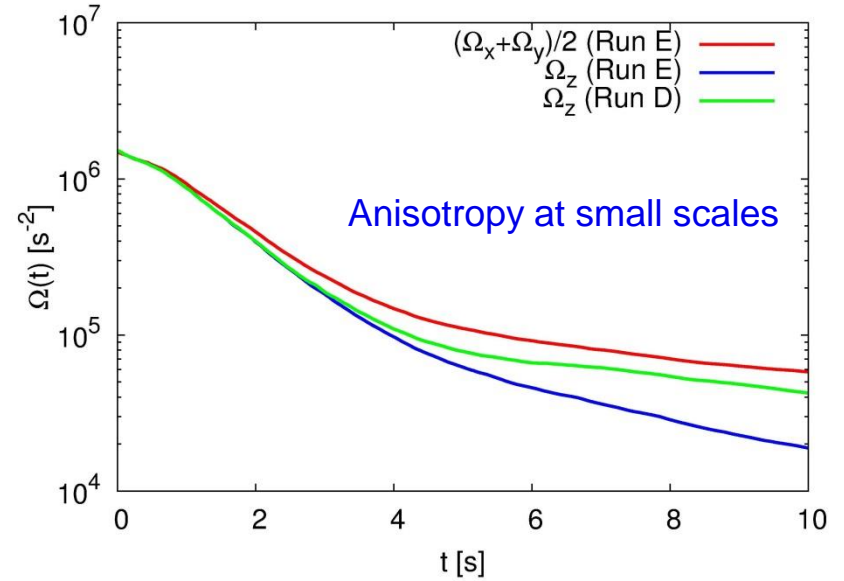
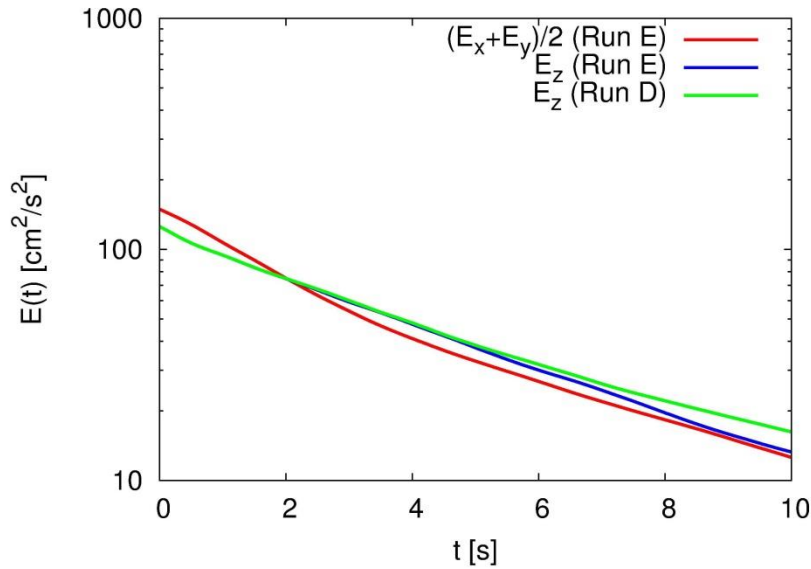


Run E

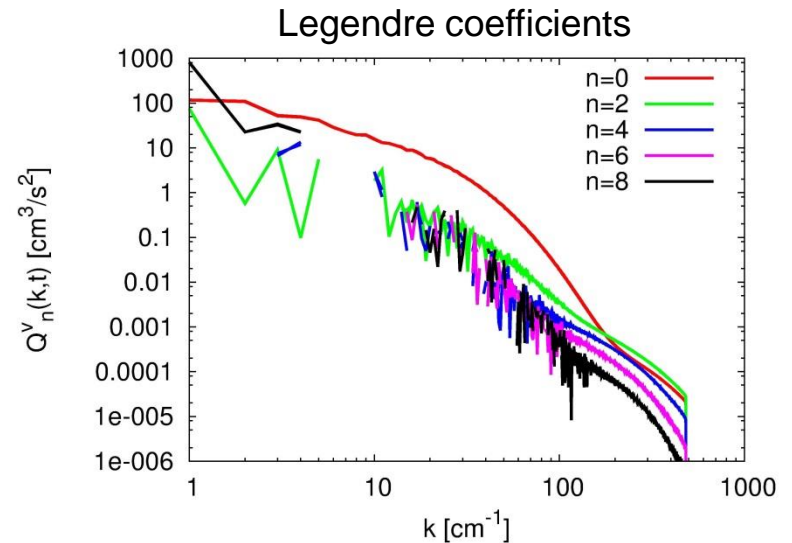
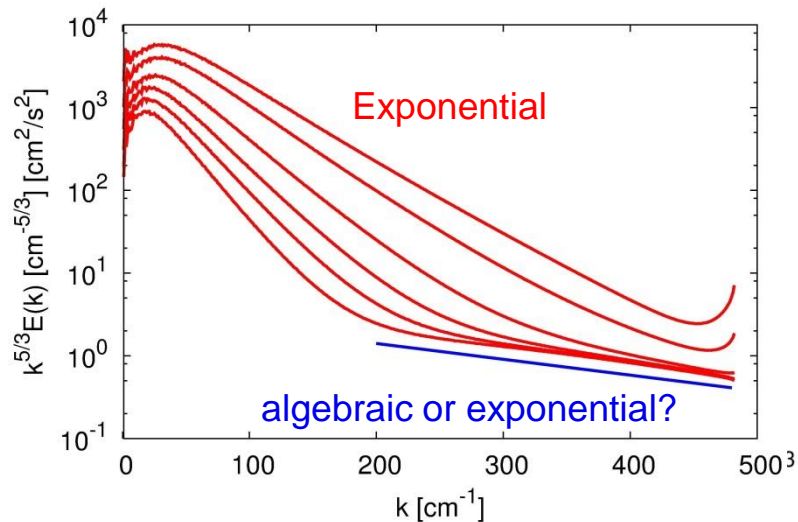


Power law $g(R) \propto R^{-\alpha}$

Turbulence modulation



Spectrum modulation begins at high wavenumbers and proceeds low wavenumbers



Turbulence spectrum

$\log E(k)$
 $\log E_\theta(k)$

Kolmogorov picture
energy cascade from large to small scales

Turbulence in cloud
Forcing at small scale
Inverse cascade?

Inertial-convective range

$k^{-5/3}$
 $\Pi \quad \Pi_\theta$

$-\hat{\Pi} ?$

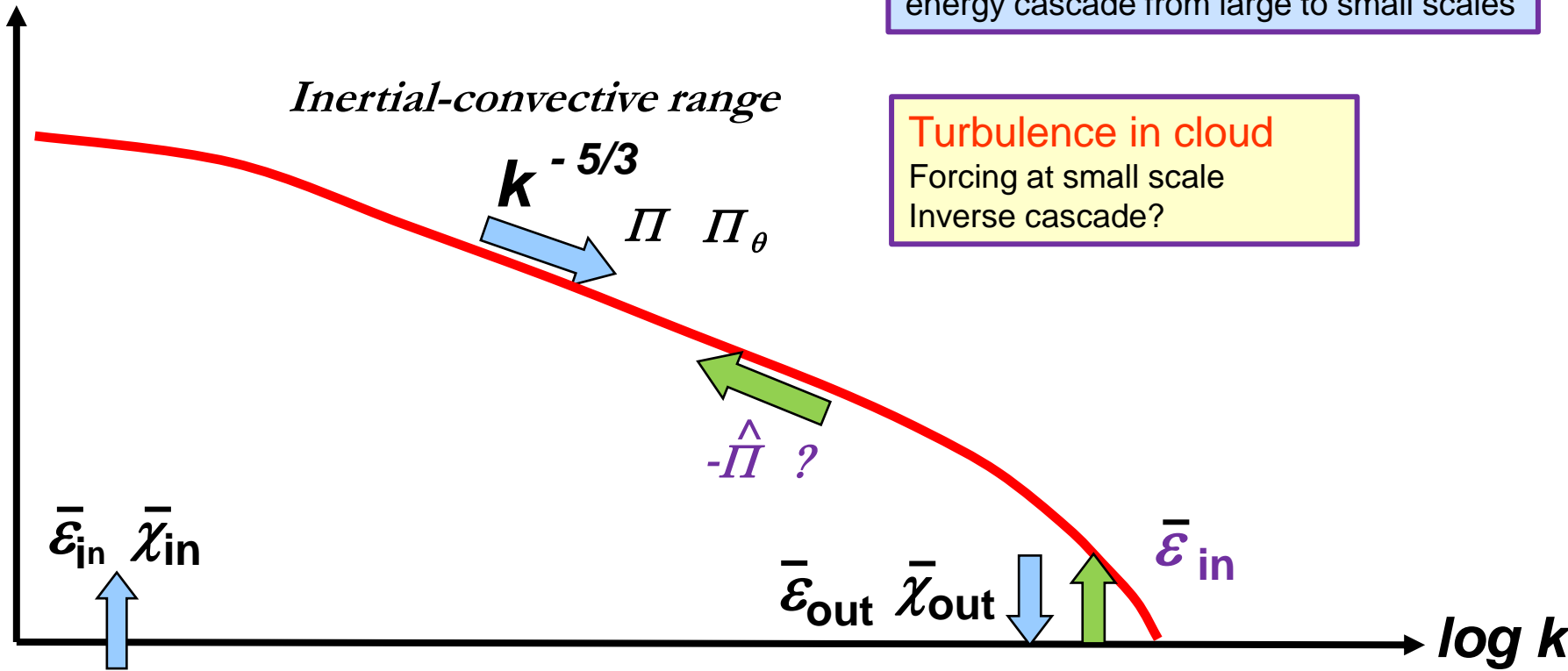
$\bar{\epsilon}_{in} \quad \bar{\chi}_{in}$

$\bar{\epsilon}_{out} \quad \bar{\chi}_{out}$

$\bar{\epsilon}_{in}$

$\log k_d$
($S_c \sim 1$)

$\log k$



Contribution of buoyancy force

$$\frac{\partial u}{\partial t} + u \cdot \nabla u = -\nabla p + \nu \nabla^2 u + e_z B + f, \quad \nabla \cdot u = 0$$

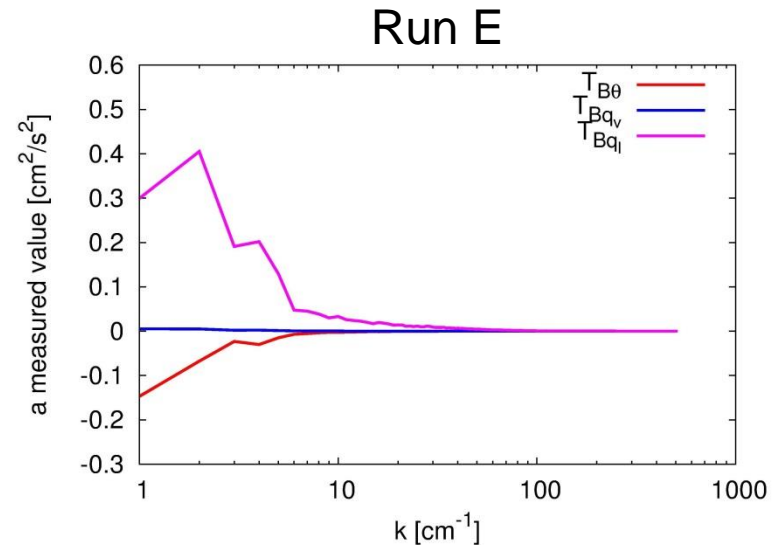
$$B = g \left(\frac{T - T_0}{T_0} + \epsilon(q_v - q_{v0}) - q_l \right) = B_\theta + B_q + B_l$$

$$\frac{DE}{Dt} = -\epsilon + P_B, \quad P_B = \langle B u_3 \rangle = \int T_B(k, t) dk$$

$$T_B(k, t) = 2 \sum_{\text{shell}} \text{Real} (\langle B_\theta(k, t) u_3(-k, t) \rangle + \langle B_q(k, t) u_3(-k, t) \rangle + \langle B_l(k, t) u_3(-k, t) \rangle)$$

$$= T_{B\theta}(k, t) + T_{Bq}(k, t) + \underline{T_{Bl}(k, t)}$$

dominant



Summary

- Reynolds number effects (no collision)
may be small on growth rate of the average particle radius
broaden PDF of particle radius
- Average growth of cloud droplets governed by diffusion process is very slow
[Needs for particle collision process](#)
- Turbulence modulation begins at small scales due to the cloud droplets
and the anisotropic modulation is inversely transferred to large scales

雲粒子間衝突

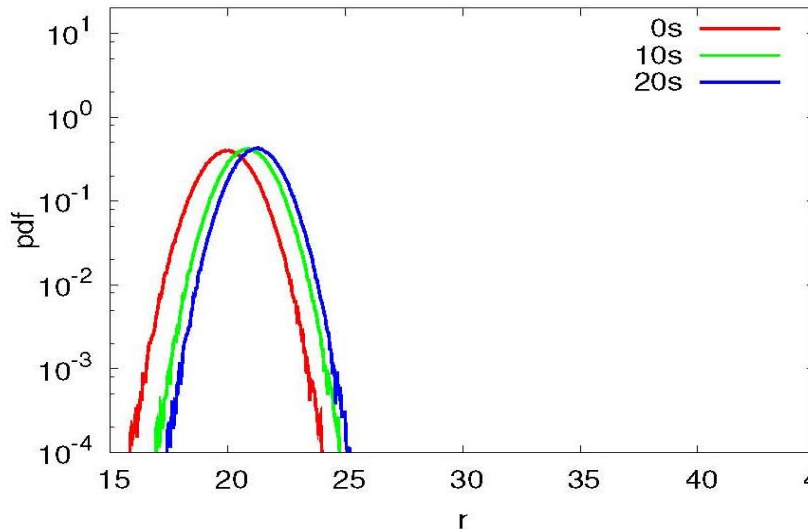
衝突条件

$$r = \sqrt{(x_i - x_j)^2 + (y_i - y_j)^2 + (z_i - z_j)^2} \leq r_i + r_j$$

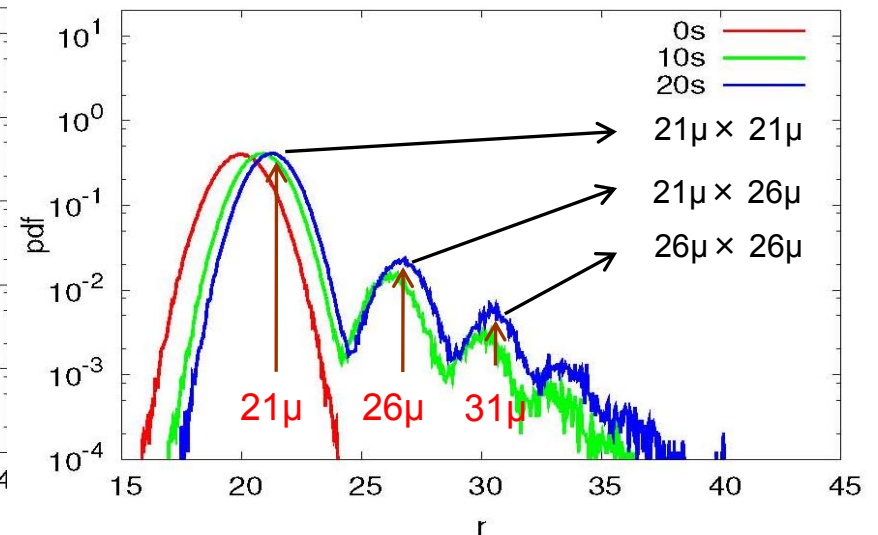
・質量保存 $M_p = M_i + M_j \quad \longrightarrow \quad r_p^3 = r_i^3 + r_j^3$

・運動量保存 $M_p \mathbf{V}_p = M_i \mathbf{V}_i + M_j \mathbf{V}_j \quad \longrightarrow \quad \mathbf{V}_p = \frac{M_i \mathbf{V}_i + M_j \mathbf{V}_j}{M_p}$

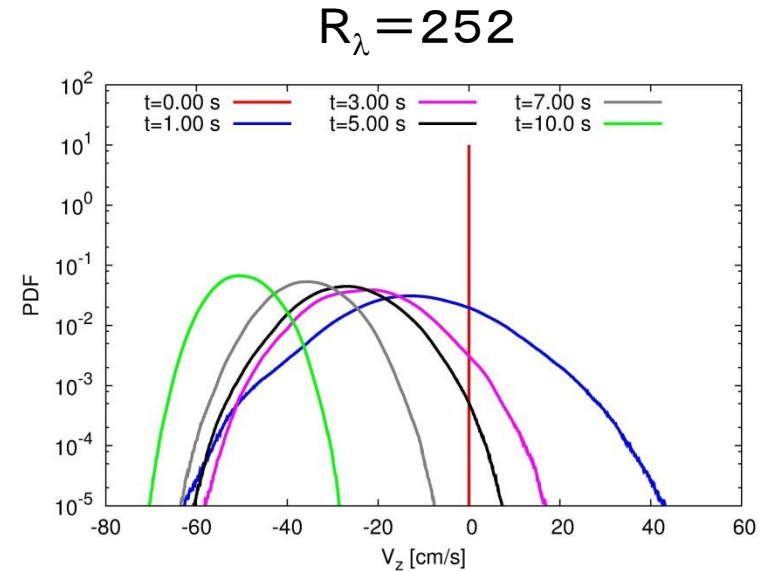
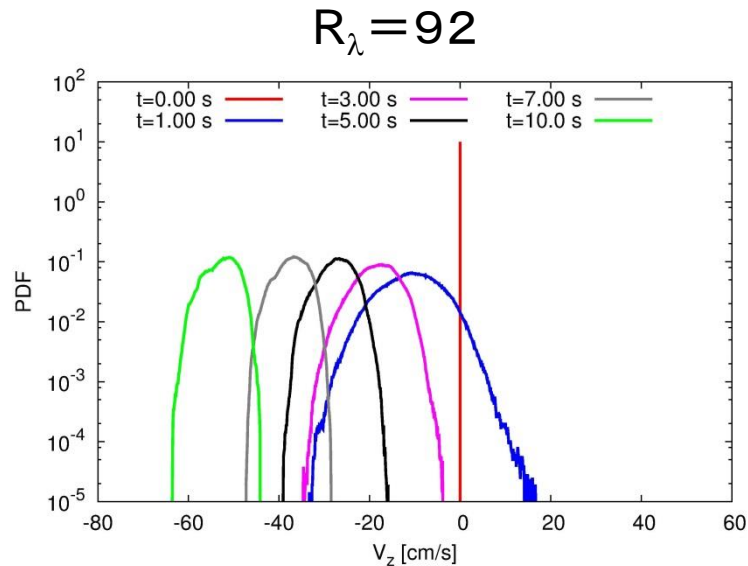
衝突・合体なし



衝突・合体あり



PDF of vertical velocity of cloud droplets



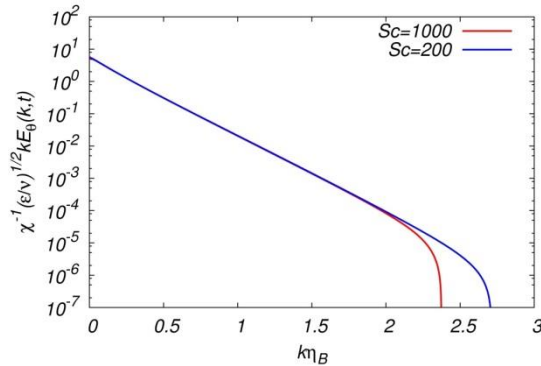
When R_λ increases

Wider PDF of u_3 \longrightarrow More chance to collide

Larger R_λ \longrightarrow Faster growth and wider PDF of r

Other progress

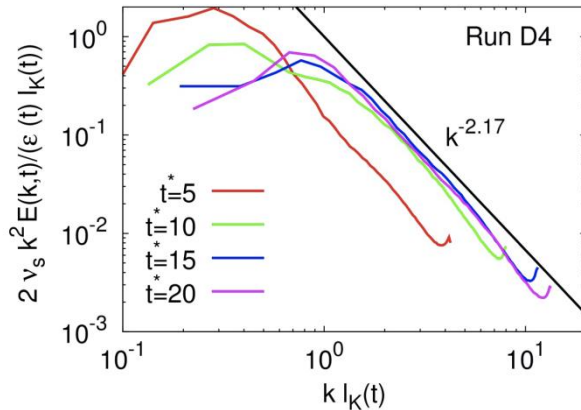
- Scalar spectrum at Sc=1000



$$E_{1\theta}(k) = C_B \chi (\epsilon/\nu)^{-1/2} k^{-1} \exp\left(-\sqrt{6C_B} k \eta_B\right)$$

Gotoh

- Polymer effects on turbulence



Power law behavior of the energy spectrum in the far dissipation range

Watanabe and Gotoh

- Fast Time-reversible method for MD of water molecules

Kajima and Ogata

Acknowledgements

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